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3 **Human origins in a southern African palaeo-wetland and first migrations**

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Anatomically modern humans (AMHs) arose in Africa some 200 thousand years ago (ka)¹⁻⁴. While some of the oldest skeletal remains suggest an eastern African origin², southern Africa is home to contemporary populations representing the earliest branch of human genetic phylogeny^{5,6}. Generating the largest resource for the poorly represented and deepest-rooting maternal L0 mitochondrial DNA branch (198 new for a total of 1,217 mitogenomes), we show geographic isolation of L0d1'2, L0k and L0g KhoeSan descendants south of the Zambezi River. Establishing mitogenomic timelines, frequencies and dispersals, we propose L0 emerged within the residual Makgadikgadi-Okavango palaeo-wetland of southern Africa⁷, some 200 ka (95% CI: 165-240 ka). Genetic divergence points to sustained 70,000 year-long existence before an out-of-homeland northeast-southwest dispersal, between 130 and 110 ka. Palaeo-climate proxy and model data suggest increased humidity opened green corridors, first to the northeast then southwest. Subsequent homeland drying corresponds with sustained effective population size (L0k), while wet-dry cycles and likely adaptation to marine foraging, allowed the southwest migrants to achieve population growth (L0d1'2), as supported by extensive south-coastal archaeological evidence⁸⁻¹⁰. Pinpointing a southern African human origin, we describe a sustained homeland occupation prior to the first human migrations, providing evidence for climate-driven dispersals.

Southern Africa has been a long-held contender for the origin of AMHs. Home to contemporary populations representing our earliest human lineages, evolutionary time estimates have largely been based on mitochondrial DNA (mitogenomes)^{1,6}. The maternal human phylogenetic tree consists of two major branches, the extensive L1'6, including the out-of-Africa ancestral L3 sub-branch (or haplogroup), and the rare deep-rooting L0. L0 is predominated by southern African haplogroups: L0d, L0k, and recently described L0g⁶. In contrast, the rare L0f and common L0a are dispersed throughout sub-Saharan Africa^{1,3,6}.

Through L0 pre-screening, we identified 198 southern Africans with poorly represented haplogroups to undergo mitogenome sequencing (Supplementary Table 1), allowing for a combined analysis of 1,217 mitogenomes (**Fig. 1a** and Extended Data Table 1).

We ethno-linguistically classified study participants as KhoeSan, southern African click-speaking foragers, or non-KhoeSan. Non-KhoeSan with KhoeSan-derived L0 mitogenomes, are referred as KhoeSan-ancestral, with further geographical classification (**Fig. 1b** and Extended Data Table 2). Contemporary KhoeSan include Kalahari (Kw'a, Tuu and central Khoe-Kwadi speakers) and West-coastal (Khoe-Kwadi Nama speakers) KhoeSan¹¹. Southern Bantu speaking peoples who migrated down the east coast of Africa some 1,500 years ago, may have acquired an East-coastal KhoeSan heritage¹². The arrival of European colonists to the Cape in mid-1600's gave rise to the South African Coloured and Namibian Baster populations, acquiring a Cape KhoeSan heritage¹³. Excluding the east African click-speaking Sandawe and Hadza, indigenous KhoeSan populations appear to be absent northeast of the Zambezi River, supported by lack of skeletal remains representing KhoeSan-like hunter-forager morphology¹⁴. We classified the 198 new mitogenomes as: Kalahari (n=18), West-coastal (n=21), Cape (n=109), and East-coastal (n=29) KhoeSan, or non-KhoeSan (Bantu, n=19), while two were unknown. Using these identifiers, we have provided a best-fit classification for all 1,217 L0 mitogenomes (Supplementary Table 2).

Phylogenetic analysis confirms major L0 haplogroups, with the exclusion of L0b (Extended Data Fig. 1). Using a subset of 461 mitogenomes, including all rare lineages, we establish within L0 coalescence times (**Fig. 2a** and Supplementary Table 3) and use the complete dataset to reconstruct geographic dispersals (**Fig. 2b**). We redefine L0 emergence ~50-25 thousand years (kyr) prior to previous estimates^{1,6}, at ~200 ka (95% CI: 165-240 ka). L0d'k (n=309 ~187 ka) is largely KhoeSan-specific, emerging ~20 kyr prior to the widely dispersed L0a'b'f'g sister-branch (n=152 ~164 ka). Although exact branch resolution for L0k

remains undetermined, we observe a preference for L0d'k (posterior probability ~0.6) over L0a'b'f'g'k (~0.4). Irrespective, L0k (n=113) appears to remain stable for ~130 kyr before diverging into the Kalahari-specific L0k1, predominated by L0k1a (85 of 94), and rarer L0k1b and L0k2 lineages distributed around the Zambezi (Extended Data Fig. 2a). L0d remains stable for ~58 kyr before splitting into the KhoeSan-specific L0d1'2 and rarer L0d3.

Coalescing ~113 ka, L0d2 (n=226) emerges ~15 kyr prior to L0d1 (n=452). Within L0d2 (~91 ka), L0d2c is the earliest diverged (n=53 ~84 ka) with broad and almost even KhoeSan-regional distribution (Extended Data Fig. 3 and Supplementary Table 4). In 2014, we derived an ancient L0d2c1c mitogenome from the skeleton of a 2,330(±25)-year-old Cape-coastal marine forager (StHe/UCT606)¹⁵. Predating archaeological evidence for sheep-herding in the region^{12,16}, we proposed this L0d2c sub-clade represent a pre-pastoral indigenous southern African lineage. Recently, whole genome sequencing confirmed a unique southern African heritage, while two younger (<2k years) Cape skeletons showed a genetic link to east Africa and associated pastoralist migration¹⁷. Observing a bias towards Kalahari representation for L0d2b (28/44 ~65 ka) and L0d2a (62/118 ~60 ka), doubling the contribution for L0d2d (6/11) we show broad KhoeSan distribution (Extended Data Fig. 3 and Supplementary Tables 5,6). As with L0d2, L0d1 is spread throughout the KhoeSan-regional identifier, with notable over-representation including: L0d1b (104/174 ~69 ka) and L0d1c (151/184 ~59 ka) within the Kalahari, and L0d1a (32/91 ~44 ka) within the Cape (Extended Data Fig. 4). We contribute two new KhoeSan-ancestral L0d1d mitogenomes to the single published⁶.

In contrast to L01'2, L0d3 is not specific to southern Africa. While L0d3b (~30 ka) appears to be KhoeSan-specific, the rarer L0d3a (~42 ka) is exclusively found north of the Zambezi. Notably, three of six L0d3a mitogenomes were derived from click-speaking east African Sandawe. Our data supports previous studies suggesting a genetic link between east Africa and the earliest southern Africans¹⁷, last sharing a common ancestor ~59 ka. Providing

the largest mitogenome contribution to L0d3 (27/40), we observe over-representation of L0d3b in the Cape KhoeSan identifier (21/34), with notable absence within the Namibian Baster population (Extended Data Fig. 2 and Supplementary Table 7). Previously suggesting a maternal KhoeSan-ancestral distinguishing identifier for the Coloured and Baster populations¹³, we show L0d3b to be specific to the Coloured and the new L0d2b1a2a sub-clade to the Baster population (Extended Data Fig. 3b).

Within L0a'b'f'g, L0f is highly divergent (~125 ka 95% CI: 101-149 ka). Adding five L0f mitogenomes, L0f1 (13/27 ~113 ka) predominates south and L0f2'3 (14/27 ~121 ka) north of the Zambezi (Extended Data Fig. 2 and Supplementary Table 8). Within L0f1 we recognise three new branches: the northeast sister-clades L0f1c (Zambian) and L0f1b (Tanzanian), and the East-to-Cape-coastal southern clade L0f1a (n=8). Lack of L0f representation within contemporary KhoeSan, suggests L0f1a presence within southern Africa is likely a result of more recent east-coastal agropastoral back-migration. While L0a'g coalesce ~117 ka (95% CI: 94-145 ka), contributing 19 southern African to 347 L0a mitogenomes, we concur that L0a likely diverged northeast of the Zambezi (~85 ka) and spread throughout Africa³, with southern L0a1b and L0a2a representation a result of Bantu back-migration (Extended Data Fig. 5). First described within a Kw'a speaking hunter-gatherer⁶, we contribute three new and reclassify five published mitogenomes as L0g (Extended Data Fig. 2 and Supplementary Table 9). With broad KhoeSan and KhoeSan-ancestral distribution, we hypothesize that L0g diverged southwest of the Zambezi (~69 ka), akin to L0d1'2.

Our data suggests that the Greater Zambezi River Basin region, particularly the southwest Kalahari, played a significant role in shaping AMH emergence and prehistory. Today a semi-desert, salt pans within northern Botswana represent desiccated vestiges of Palaeo-Lake Makgadikgadi, which at its peak would have been the largest lake in Africa, flanked by smaller lakes from the Upper Zambezi to the Kafue Rivers^{7,18}. Contraction of the

lake, accompanied by development of the Okavango Delta, as a result of neotectonic rifting, would have created a residual wetland favourable for human and broad mammal habitation¹⁹ (**Fig. 2c**). Today the harsh Kalahari climate and oxygen-rich salt pans are not ideal for fossil and pollen preservation, respectively. However, lithic artefacts representing the Middle to Late Stone Age, are documented from the Makgadikgadi Pans and surroundings^{7,20,21}, while palynology suggests a once grassland and forest biome²². Our data further suggests the Makgadikgadi-Okavango palaeo-wetland sustained AMH existence for ~70 kyr, supported by mitochondrial data for ancestral giraffe, lion and zebra^{23–25}, before out-of-homeland migrations split founder homeland populations L0d, L0f and L0a'g.

Southwest of the homeland L0d1'2 experienced episodic splits, with broad south-coastal occupation, while L0g is less successful. Bayesian Skyline Plot (BSP, **Fig. 2c insets**) analysis confirms effective population growth for L0d1'2, while extensive archaeological evidence suggests cognitive human behaviour at the southern tip of Africa between ~100-60 ka^{8–10} with associated increase in the density of coastal over inland time-appropriate archaeological sites²⁶. Northeast of the homeland, L0d3 and L0f are less successful, while L0a underwent significant population trajectory post-dating the out-of-Africa migration (BSP L0a). The northeast migration route is supported by the appearance of data-appropriate archaeological sites²⁶. Within the homeland, L0k sustained a constant effective population size (BSP L0k), as did Kalahari-predominant L0d2b, L0d2a and L0d1c. While the presence of L0k in Zambia has been suggested to represent contact with an ancient pre-Bantu population²⁷, we propose these rare lineages represent an ancient out-of-homeland KhoeSan offshoot.

Orbitally-driven large-scale hydroclimate variations have been proposed as a pacemaker for early human migrations^{28,29}, with wetter conditions and resulting 'green corridors' invoked to explain the out-of-Africa migration ("pull") or drier conditions and resulting food shortages forcing dispersals ("push")³⁰. To determine whether our predicted

homeland isolation and major dispersals may have been driven by climate shifts, we analysed four key palaeo-hydroclimate datasets^{29,31–33}, along with a transient 784 kyr-long glacial/interglacial simulation conducted with the LOVECLIM earth system model²⁸ (**Fig. 3**). Although limited by available palaeo-proxy records and a climate model of intermediate complexity, we observed a considerable degree of coherence on orbital timescales (Extended Data Fig. 6). During the homeland period (200–130 ka), palaeo-data links the 21,000 year-long precession cycle (**Fig. 3a**), with three wet-dry cycles (**Fig. 3b**). Conversely, climate model simulates an extended drought, owing to a more pronounced eccentricity signal (**Fig. 3e**), proposing a wetland oasis in an otherwise vast harsh environment.

During the out-of-homeland period (130–110 ka), our model simulation supports humid conditions to the northeast facilitated the first dispersals, concurring with L0f coalescence (~125 ka) (**Fig. 3d**). In contrast, the region southwest of the homeland experienced a ~15 kyr-long megadrought before an orbital shift created favourable humid conditions for L0d1'2 dispersal (~113 ka) (**Fig. 3f**), as supported by palaeo-data (**Fig. 3c**). This is also around the time the northeast L0a and southwest L0g migrants last share a common ancestor (~117 ka). During the last glacial period (~100–11 ka), we observe reduced amplitude of orbital-scale hydroclimate changes and overall drying within the homeland (**Fig. 3b**), while the southwest coastal hydroclimate was dominated by precessional variability and relatively agreeable environmental conditions (**Fig. 3c,f**). Interestingly, periods of slowdown and acceleration in effective population size estimates for L0d1'2, coincide with regional changes in hydroclimate, further linking climate, population size and evolution.

We propose the Makgadikgadi-Okavango palaeo-wetland as the possible homeland for AMHs. Although one cannot exclude the possibility of a polycentric origin³⁴, this deltaic-lacustrine ecosystem would have provided an ideal geographic locality for the evolution and 70 kyr sustained existence for AMHs deepest-branching maternal founder population. Increased

humid conditions, supported by palaeo-lake system reconstructions³⁵, between 130 and 110 ka would have opened green corridors for successful northeast-southwest migrations, supporting a “pull” scenario. Drying within the homeland following the out-of-homeland period, supported by hydroclimate data (110-100 ka) and model simulation (100-80 ka), would have created a “push” scenario, where reduced land carrying capacity would have increased pressure to seek out climatically more favourable regions. We speculate, the southwest migrants maintain a successful coastal forager existence, while the northeast migrants, as with the later branching L1’6, gave rise to ancestral pastoral and farming populations. A recent publication, provides further mitochondrial evidence to support the northeast out-of-homeland migration route and expansion into east Africa around 70-60 ka³⁶. Revealing a southern African homeland for L0 emergence and extended subsistence, we propose an out-of-homeland migration event, likely driven by astronomically-induced regional shifts in hydroclimate, shaped present-day ethnic/genetic diversity.

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Author Information Reprints and permissions information is available at www.nature.com/reprints. Current affiliation for D.C.P. is The Centre for Proteomic and Genomic Research, Cape Town, South Africa and for H.R. is Akademia, Johannesburg, South Africa. The authors declare no competing financial interests. Readers are welcome to comment on the online version of the paper. Correspondence should be addressed to V.M.H.

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Figure Legends

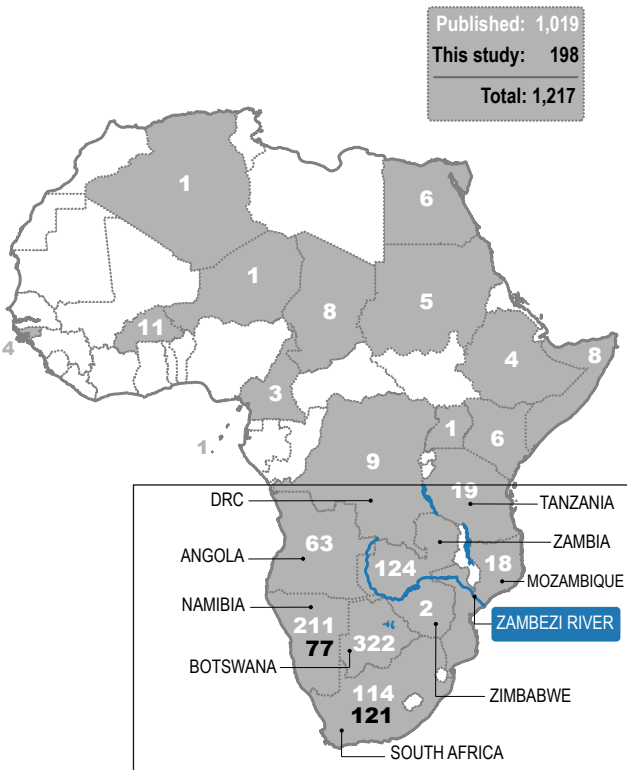
Fig. 1. Geographical distribution of 1,217 L0-mitogenomes. **a**, Countries within (n=1,139) or outside (n=78) Africa from where L0-mitogenomes were sourced, including 198 new L0-mitogenomes (black numbers). **b**, Present day southern Africa with geographical distribution of KhoeSan population identifiers defined as; KhoeSan (orange): Kalahari and West-coastal, or KhoeSan-ancestral (green): Cape or East-coastal. The Zambezi River provides a geographic divide between the KhoeSan and largely non-KhoeSan population identifier.

Fig. 2. L0 Phylogenetic tree, major southern African L0-haplogroup geographic distributions and out-of-homeland L0 dispersal routes. **a**, Phylogenetic branching and coalescence times derived from a focused subset of 461 L0-mitogenomes, including all rare branches, and anchored to Neanderthal (n=7). The Somalian-derived (Som20) L0d3 mitogenome³ could not be assigned. **b**, Geographic distribution (identifiers described in Fig. 1b) for all KhoeSan-specific mitogenomes (of 1,217): L0d3 (n=40), L0d1'2 (n=677, excluding one unknown), L0k (n=105, excluding seven L0k1b and a single Yemen-derived L0k2), and L0f1 (n=13). Predominant geographic representation (shaded regions), with region specific overflow represented by total number of mitogenomes, including country specific representation north of the Zambezi. **c**, Schematic map of southern Africa representing the Makgadikgadi-Okavango palaeo-wetland sustained AMH homeland (200-130 ka), supported by archeological⁷ and genetic wildlife data²³⁻²⁵. The out-of-homeland migration (130-110 ka), results in L0d, L0a'g and L0f splitting. L0d3, L0a and L0f migrate in a northeast direction, L0d1'2 and L0g migrate southwest, while L0k remains in the homeland. Bayesian Skyline Plot

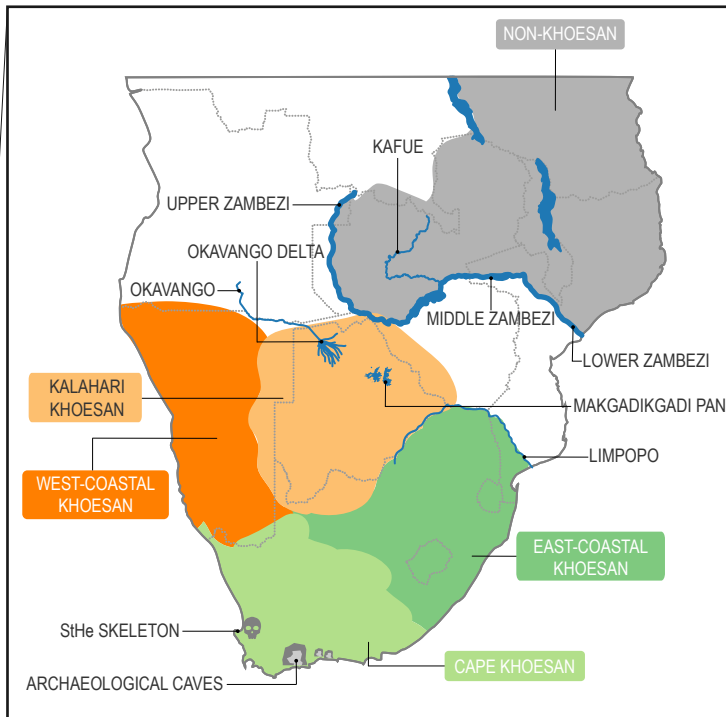
(BSP) analyses [insets] for major L0-haplogroups over time predicts: maintenance of the homeland L0k population (orange), population growth for the broadly dispersed southwest L0d1'2 migrants (purple), supported by archaeological evidence (100-60 ka)⁸⁻¹⁰ and StHe mitogenome¹⁵, while population growth for the northeast L0a migrants coincides with the out-of-Africa migration (aqua).

Fig. 3. Reconstructed (left panel) and simulated (right panel) climatic conditions during the out-of-homeland migration. **a**, Austral summer insolation changes (blue) at 27°S. **b**, Eastern and central southern Africa hydroclimate composite (shading) obtained by averaging Fe/K runoff record from core CD154-1006P³¹ and the Pretoria Salt Pan rainfall reconstruction²⁹, extended from 250-190 ka (grey line), and depicting Bayesian Skyline Plot (BSP) effective population size for homeland L0k (orange dashed). **c**, Southwestern hydroclimate reconstruction (shading) obtained by averaging leaf wax data (MD08-3167³²) and aridity index from cores (MD96-2094³³), with aridity record extended from 250-140 ka (grey line) and L0d1'2 BSP effective population size (purple dashed). **d**, Simulated LOVECLIM normalized precipitation changes (shading) northeast of the homeland (33°E, 13°S) and coalescence time probabilities for L0f sub-branches (blue bell curves). **e**, Same as for d, but for the homeland and coalescence probabilities for L0, L0d'k, L0a'b'f'g (black) and L0k sub-branches (orange). **f**, Same as d, but for the area southwest of the homeland (17°E, 30°S) and L0d1'2 coalescence times (purple). Blue bars indicate predicted Makgadikgadi high stand phases³⁵.

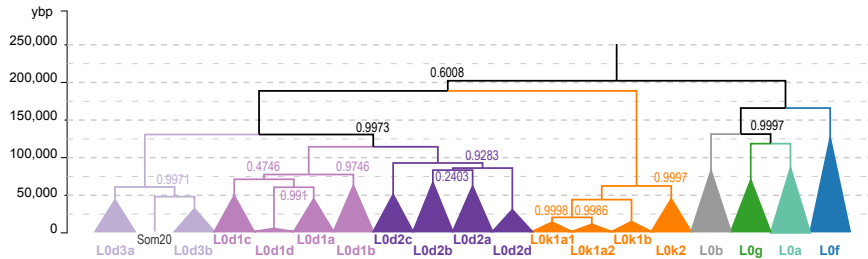
a. Geographic Distribution of L0 Mitogenomes



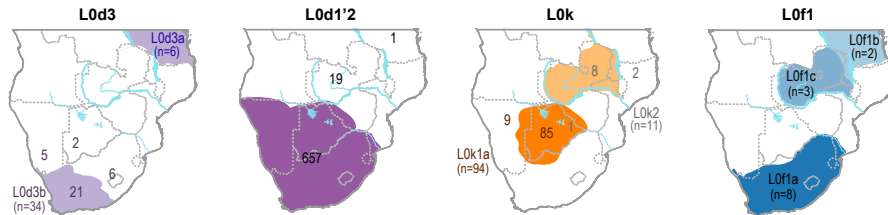
b. Geographic Distribution of Khoesan Identifier



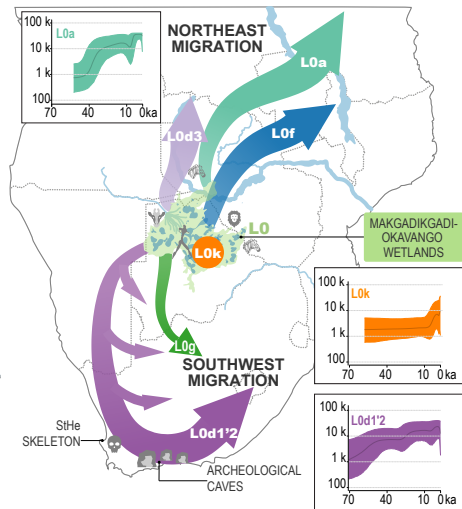
a. Phylogenetic Tree For 461 L0 Mitogenomes

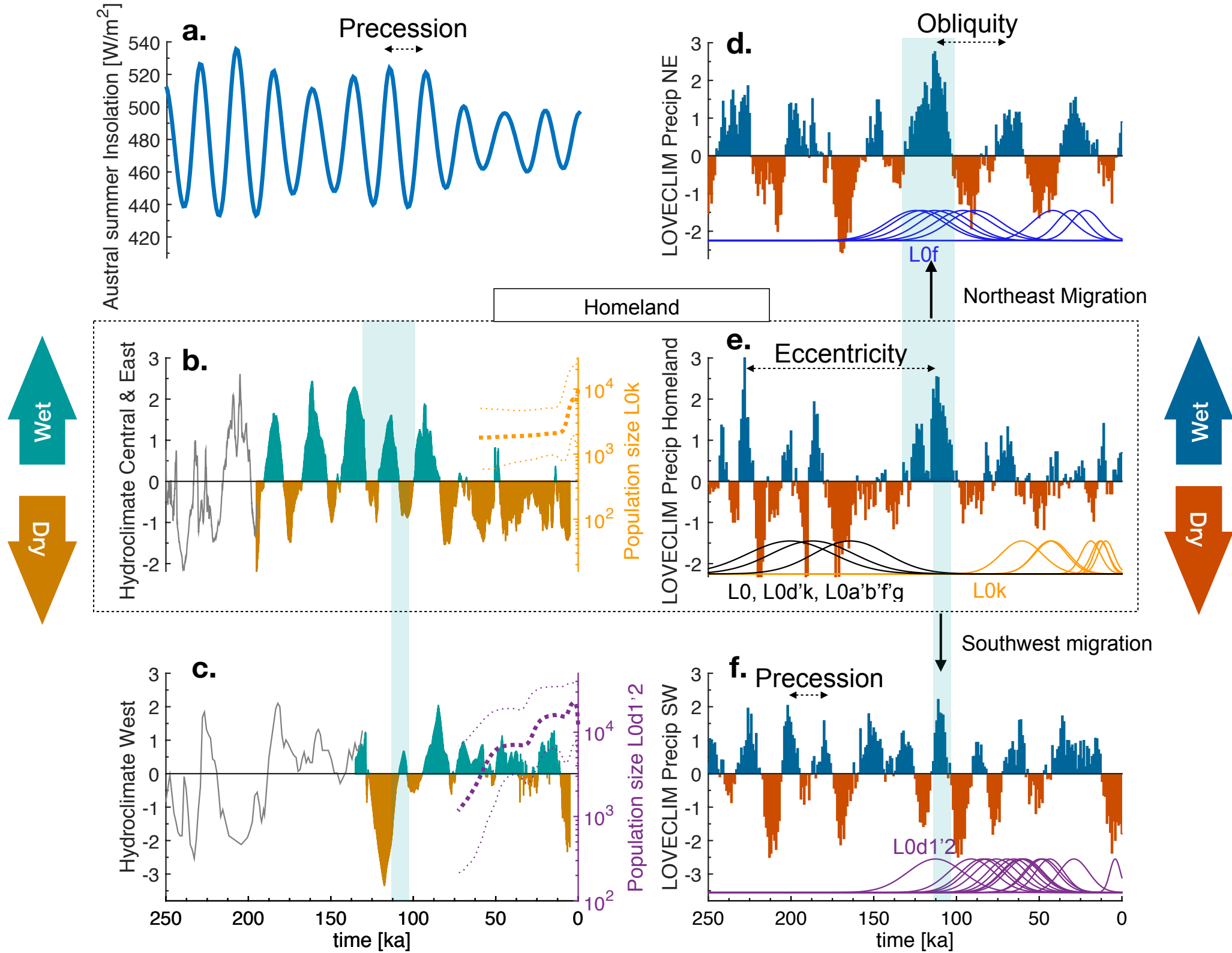


b. Geographic Distribution of L0d, L0k, and L0f from 1,217 L0 Mitogenomes



c. Migrations





METHODS

Statement on population identifiers. The authors acknowledge that population identifiers (or ethnic labels) have different meanings for different peoples across different countries and between and within different ethnic groups. During the apartheid rule, South Africans were grouped according to ethnic identities, which resulted in discrimination based on population identifiers such as Bantu or Coloured. In turn, others view the very same population identifiers with cultural identity and pride. In 2013, we performed a study led by a Coloured co-author to assess the sensitivity in self-identification as Coloured. Of 521 participants, 91.2% self-identified as Coloured, Cape Coloured or South African Coloured, while 8.8% elected against the use of Coloured for self-identification¹⁴. In turn, using such population identifiers within the context of the United States would be seen as derogatory and highly offensive. We have previously genetically profiled the Baster population of Namibia¹³ and again what could be to others a derogatory term, to the Baster community of Rehoboth in Namibia, the term is used with immense pride, with recognizing themselves as a Republic with a national flag³⁷.

In this study, the authors have used linguistics, supported by ethnicity, to provide population identification, with further historical, geographic and genetic classification for deriving maternal contributions (described in the next section). KhoeSan (or KhoeSaan) languages are grouped together due to their use of click consonants as a unique language identifier. Once spread over the entire southern African region, KhoeSan languages are today restricted largely to populations residing in Namibia and Botswana (and southern Angola), while two Tanzanian isolates, Sandawe and Hadza, are believed to be linguistically-related click-languages (or East African KhoeSan)³⁸. Meaning literally ‘forager’ (San) ‘person’ (Khoe), culturally the KhoiSan identifier refers to forager hunters (San) or herders (Khoi). At times linguistic and cultural identities clash. For example, Nama and Hai/om peoples both speak the Khoe-Kwadi based language Nama, while culturally and historically these two

389 populations are quite different, representing a herder and hunter-gatherer ancestry,
390 respectively. Additionally, autosomal genetic data has been used to provide further insights
391 into KhoeSan admixture and substructures, highlighting, for example at a genetic-level
392 historical differences between the Nama and Hai//om³⁹. The authors have attempted to capture
393 both ethnic and linguistic identifiers that best reflect population ancestry. In contrast to
394 KhoeSan languages, Bantu languages are not traditionally click-based, yet again exceptions
395 exist within southern African Bantu languages, for example isiXhosa and isiZulu languages
396 have borrowed click consonants from their KhoeSan neighbours. Spoken across the entire sub-
397 Saharan Africa (up to 500 groups), the Guthrie classification of languages further identifies the
398 S-zone or Southern Bantu (South Africa, Zimbabwe, southern Mozambique and most of
399 Botswana) and the R-zone or Southwest Bantu languages (northern Namibia, southern Angola
400 and northwest Botswana)⁴⁰, which are of relevance to this study.

401 **Ethics statement and recruitment.** The study was performed in accordance with the ethical
402 standards of the overseeing human research ethics committees and local governance, as per the
403 1964 Helsinki Declaration. The study was reviewed and approved by the Ministry of Health
404 and Social Services (MoHSS) in Namibia (#17-3-3 2008, 2014 and 2019), with additional local
405 approvals from community leaders, the University of Pretoria Human Research Ethics
406 Committee (HREC #43/2010 and HREC #280/2017), including US Federal-wide assurance
407 (FWA00002567 and IRB00002235 IORG0001762), as well as the South African National
408 Blood Service (SANBS) HREC (HREC #2012/11). Participants were recruited within the
409 borders of Namibia and South Africa, with self-reported ethno-linguistic population identifiers
410 recorded. Blood samples were taken after receiving written and/or recorded informed consent.
411 Isolated DNA was shipped under the Republic of South Africa Department of Health Export
412 Permit (#J1/2/4/2), in accordance with the National Health Act 2003, to the Garvan Institute of

Medical Research in Australia. Mitogenome sequencing was performed in accordance with site-specific approval granted by St Vincent's Hospital HREC in Australia (SVH 15/227).

Participant population identifiers. Merging with published data for a total of 1,217 L0 mitogenomes, participants were broadly classified as KhoeSan, Bantu or Cape multi-ethnic heritage. Indigenous KhoeSan inhabiting the inland semi-desert Kalahari region of Botswana and Namibia include the Kw'a (Ju/'hoan or Hoan, and !Xun or !Xuun), Tuu (or Taa) and Khoe-Kwadi (Naro, //Ani, Khwe, Buga, G//ana, G//ui, //Xokhoe, Tshwa and Shua) click-speakers. Indigenous KhoeSan inhabiting the west coastal region of Namibia speak a Khoe-Kwadi or Nama language and include the Nama, Damara, Topnaar (#Aonin) and Hai//om click-speakers^{41,42}. Novel mitogenomes were derived from 15 Kalahari KhoeSan, including Ju/'hoan (n=9), !Xun (n=1) and Naro (n=5), and 21 West-coastal KhoeSan, including Nama (n=7), Damara (n=8) and Topnaar (#Aonin, n=6) from Namibia. Southwest Bantu (non-KhoeSan) non-click speakers of Namibia, Botswana and southerly borders of Angola, presenting with KhoeSan-predominant L0 maternal lineages, most likely carry a Kalahari or West-coastal KhoeSan mitogenome. As a result of refuge provided to the Herero by the Kalahari KhoeSan during the early 1900 German South West African genocide⁴³, we speculate in this study a probable Kalahari KhoeSan heritage for the three Herero mitogenomes.

While indigenous KhoeSan are arguably absent from the coastal regions of South Africa, KhoeSan skeletal remains spread across the region⁴⁴. Hunter-gatherer KhoeSan once inhabited a broad southwest to east-coastal region at the tip of Africa. These skeletal remains predate archaeological evidence supporting the arrival of sheep-herders who appear to have crossed the Okavango River in northern Namibia ~2.2 kya, migrating along the southwest coast to the southern Cape by ~2 kya^{12,16,20,44}. Recently, Cape KhoeSan skeletons younger than 2 ka have been genetically linked to east Africa and herder migration¹⁷. Migrating herders may have acquired indigenous KhoeSan maternal contribution. Along the east coast, southward

438 migrating non-click-speaking Bantu farmers (Southern Bantu) enter South Africa ~1,500 ya,
 439 while a second wave of Bantu migrants (Southwest Bantu) cross central Africa into Namibia
 440 ~800 ya¹². Maternal contributions to the South African Southern Bantu speaking populations
 441 (n=43, this study) may therefore either be of Bantu origin (in this case L0a lineages and
 442 therefore non-KhoeSan) or of east-coastal KhoeSan-ancestry. The arrival of European colonists
 443 and Dutch-east Indian slaves to the Cape in the mid-1600's, gave rise to multi-ethnic
 444 (European, Asian, KhoeSan and Bantu) Cape population, the ancestors of the South African
 445 Coloured (n=90, this study) and Namibian Basters (n=24, this study), speaking historically a
 446 Dutch-derived language, known as Afrikaans^{13,45}. Emerging from a common historical
 447 background to the Coloured, the Baster population have since the late 1800's distinguished
 448 themselves as independent from the Coloured, migrating to the Baster nation of Rehoboth in
 449 Namibia⁴⁶. While the vast majority of L0-mitogenomes represented in the Baster and Coloured
 450 population are of Cape KhoeSan heritage (100% and 94.4%, respectively), we observe a
 451 percentage of non-KhoeSan (Bantu) L0a lineages within the Coloured.

452 **L0-Haplogroup pre-screening.** Subjects were selected for whole mitogenome sequencing
 453 based on pre-screening for specific L0 markers using direct amplicon-specific Sanger
 454 sequencing. Specifically, a 2,673 bp region (rCRS position 3322–5995) was amplified and
 455 initially screen for the L0 variant T5442C. L0 samples were further screened to delineate into
 456 L0d (T4232C), L0d1 (G3438A), L0d1b (T3618C), L0d1c (C4197T), L0d1'2 (A3756G), L0d2
 457 (A3981G, C205T, A4044G), L0d2a (A5153G), L0d2d (G5147A, G5231A), L0d2C (A4038g,
 458 T4937C) and L0d3 (G5460A, G5773A). This identified a 188 samples carrying a rare L0-
 459 haplogroup: L0d1b (n=21), L0d1c (n=13), L0d2a (n=30), L0d2b (n=7), L0d2c (n=15), L0d2d
 460 (n=6), L0d3 (n=29), L0a1 (n=6), L0a2 (n=6), L0f (n=5) and L0k (n=5); as well as 55 samples
 461 that could not be unambiguously assigned to major L0 sub-lineage: L0d1a'c (n=2) L0a'b'f'k
 462 (n=5), L0a'b (n=2), L0d2 (n=1), and L0d1 (n=45, assumed L0d1a) (Supplementary Table 1).

Whole mitogenome sequencing. Mitogenomes were isolated using two overlapping amplicons as previously described^{6,47}. Specifically, two primer pairs to isolate and amplify fragments 12,250–3,005 (7.2 kb) and 2,583–12,337 (9.7 kb) of the circular mitogenome. This pair of primers has been demonstrated to effectively capture of the mitogenome with high specificity whilst minimizing off-target capture of nuclear copies of mitochondrial-derived DNA (NuMTs). Following touchdown long-range amplification with the Platinum™ Taq DNA Polymerase High Fidelity (Invitrogen), the two amplicons were purified using the AMPure XP beads (Agencourt) and combined in a 7:13 ratio of short to long fragments. Sequencing was performed on the Ion Torrent PGM platform. In brief, 200 bp single-end sequencing libraries were prepared using the Ion Xpress Plus Fragment Kit and Ion Xpress™ Barcode Adaptors (ThermoFisher), and 4 to 16 samples (barcodes) were pooled and sequenced on 314v2 Ion Chips. Using the Ion Torrent suite v5.0.2.1, sequencing reads were quality trimmed, and aligned to the human mitochondrial revised Cambridge Reference Sequence (rCRS; accession NC_012920.1). Consensus mitogenome sequences were derived by first identifying variants relative to rCRS, using samtools (v1.3.1) mpileup (with parameters -d 10000 -L 1000 -Q 7 -h 50 -o 10 -e 17 -m 4)⁴⁸ and bcftools (v1.3.1) call (with parameters -c -M) (<http://www.htslib.org/doc/bcftools.html>), then converting to fasta format using samtools' vcfutils.pl vcf2fq program.

Data availability. The consensus sequences for this set of 198 mitogenomes have been deposited to NCBI with Accession Numbers MK248274-MK248471.

Public data. An exhaustive search for publicly available L0-mitogenomes was performed between 2015 and 2017, identifying 26 studies containing a total of 6,334 mitogenomes. L0 status for all mitogenomes was deduced, either directly from the original publication or by downloading the nucleotide sequences from NCBI and evaluating their haplogroup using HaploGrep2 (v2.1.13)⁴⁹ based on PhyloTree Build 17⁵⁰. From this, a subset of 1,019 L0-

mitogenomes were identified and included in this study (Extended Data Table 1, Supplemental Table 2). Public genomes were broadly classified as KhoeSan, Bantu (KhoeSan-ancestral), or Non-KhoeSan based on the reported population and/or country of origin.

Whole mitogenome haplotyping. HaploGrep2⁴⁹ was used to type all 1,217 sequences against PhyloTree Build 17⁵⁰. This resulted in the refinement and reclassification of our 198 mitogenomes, resulting in L0d1 (n=81, including 45 L0d1a, 21 L0d1b, 13 L0d1c, and 2 L0d1d), L0d2 (n=58, including 30 L0d2a, 8 L0d2b, 14 L0d2c and 6 L0d2d), L0d3 (n=27), L0a (n=19), L0f (n=5), L0k (n=5) and L0g (n=3) mitogenomes (Supplementary Table 1). This refined, and in some cases reclassified, the haplogroups of the 1,019 public mitogenomes (Supplementary Table 2).

Phylogenetic inference. Multiple sequence alignment was performed across all 1,217 mitogenomes along with seven Neanderthal genomes (Supplemental Table 10), using MUSCLE v3.8.31⁵¹ with parameters -maxiters 3 -diags1. Phylogenetic inference was performed using FastTree v2.1.7 (SSE3)⁵² using the generalised time reversible (-gtr) and discrete gamma model with 20 rate categories (-gamma). A summary of the inferred phylogenetic tree is shown in Extended Data Fig. 1, with the tree re-rooted to the seven Neanderthal genomes.

Bayesian phylogenetic inferences and divergence times were calculated using BEAST2 v2.4.2 with BEAGLE 2.0⁵³. Due to the computational burden of this analysis, BEAST was performed on a subset of 461 mitogenomes, selected to include: (i) Only complete mitogenomes (27 mitogenomes with only the coding region^{54,55} were excluded); (ii) All 198 novel mitogenomes from this study; (iii) All 121 L0-mitogenomes from our previous studies, Chan *et al.* 2015⁶ (n=77), Morris *et al.* 2014¹⁵ (StHe, defining the new haplogroup L0d2c1c), McCrow *et al.* 2016⁴⁷ (n=37), Schuster *et al.* 2010⁵⁶ (n=6); (iv) All rare haplogroups, namely L0g (n=9), L0f (n=22), L0d3 (n=30), L0d1d (n=3), L0d2d (n=11), and L0k2 (n=12); (v) All

mitogenomes that could be unambiguously typed by HaploGrep2⁴⁹ (n=14; none from this study); and (vi) A random subset of mitogenomes for all remaining sub-lineages not already represented.

Multiple sequence alignment of the subset of 461 AMH and seven Neanderthal mitogenomes was converted to NEXUS format using the convert function of seqmagick v0.6.1 (<https://fhcrc.github.io/seqmagick>) with parameter --alphabet dna-ambiguous. This provided the input to BEAST2. Specifically, BEAUTi v2.4.2 was used to set up the phylogenetic model, assuming: (i) the Gamma Site Model with 6 gamma categories and no invariant sites; (ii) the generalized time reversible substitution model; (iii) a strict constant clock model with a normal prior of with $\mu = 1.665 \times 10^{-8}$ and $\sigma = 1.479 \times 10^{-9}$ based on Soares *et al.* 2009⁵⁷; and (iv) a Coalescent Constant Population. Times were calibrated on the seven *Homo neanderthalensis* mitogenomes with tip dates set to their reported archeological dating estimates (in years before present): Feldhofer1=40000, Vindija=38000, ElSidron=39000, Feldhofer2=40000, Mezmaiskaya=65000⁵⁸, Croatia=38310⁵⁸, Altai=50000⁵⁹ (see Supplementary Table 10). No prior was set on the most recent common ancestor of this taxon set, and calibration was applied to the leaves instead of the most recent common ancestor. Further, a normal prior, $N(\mu = 200000, \sigma = 50000)$, was set on the coalescent time of the AMH genomes, and a tip date of 2,330 ybp was set for StHe¹⁵.

Five BEAST replicates were performed, each with 100 million MCMC iterations, sampling every 10,000. Tracer v1.6 was used to evaluate BEAST trace files (Supplementary Table 11), ensuring all runs had converged. The five replicates were combined using LogCombiner v2.4.2, discarding 10% of the samples as burn-in for each replicate and without re-sampling states at a lower frequency.

Sampled trees from BEAST were summarized into a single Maximum Clade Credibility target tree using TreeAnnotator v2.4.2 for each of the five replicates, discarding the

first 10% as burn-in. To summarize across replicates, sampled trees from the five replicates were first combined using LogCombiner v2.4.2, again discarding the first 10% as burn-in from each replicate, but re-sampling at a lower frequency of 50,000 (five replicates of 10,000 samples). The combined, re-sampled trees were then summarized with TreeAnnotator v2.4.2 as for the individual replicate BEAST results.

FigTree v1.4.2+ (Rambaut 2012; <http://tree.bio.ed.ac.uk/software/figtree/>) was used to visualize all resulting trees.

Bayesian Skyline Plot analysis. BSP analyses were performed to estimate the demographic history of each maternal haplogroup. While maternal haplogroups do not necessarily equate to population data, it has been suggested that the signal associated with a haplogroup can still provide insight into the demographic processes in the populations carrying it^{60,61}.

For each haplogroup of interest (e.g. L0a, L0d1'2, and L0k), a nexus file was derived using SeqMagic v0.6.1 as described above. BSP analyses were performed using BEAST2, using BEAUTi 2 for model setup as before, with the following key differences: (i) the gamma shape of the Gamma Site Model was estimated with an exponential prior with mean = 1.0 and offset = 0.0; (ii) the molecular clock was fixed (not estimated) at 1.665×10^{-8} based on Sores *et al.*⁵⁷; and (iii) the phylogenetic tree prior was set to Coalescent Bayesian Skyline, assuming 20 intervals between the root of the tree and the present time.

Tracer v1.6 was used to reconstruct the Bayesian Skyline from the sampled trees for each analysis, using a stepwise constant variant and the lower 95% highest posterior density of the root height as the maximum time. Results of this analysis are summarized in Supplementary Table 12.

Geographic history of the palaeo-wetland Makgadigadi. Initiated around 2 million years ago, palaeo-lake Makgadikgadi originally covered an area of $\sim 170,000 \text{ km}^2$ at its highest lake

stand, bounded by a shoreline of ~995 m⁷. A degraded sand ridge (the Deception Ridge), was associated with the 995 m shore in the southwest of the lake. This lake would have covered more than twice the area of modern Lake Victoria, and like the latter, would have caused a significant climatic feedback, with locally enhanced rainfall. We previously proposed that this was, in turn, responsible for initiation of the surrounding (now-fossil) drainages, creating a well-watered environment and very favorable habitat for mammals, including hominids⁷. Smaller lakes, now represented by residual wetlands, also formed on the upper Zambezi and the modern Kafue Flats on the Kafue River, resulting in an archipelago of palaeo-lakes in south-central Africa during the early and Middle Pleistocene.

Palaeo-Makgadikgadi bounded the 995 m shoreline was originally sustained by a major drainage line, which included the Chambeshi as headwaters, connected to the upper Zambezi via the Upper Kafue River. Severance of original links between the Chambeshi River and upper Kafue, and the latter and the Upper Zambezi resulted in a sequential contraction of the Makgadikgadi to a much smaller water body. This is reflected in a series of fossil shorelines, associated with breaks in slope, at progressively lower levels (945 m, 936 m and 922 m). The Gidikwe Ridge was associated with the 945 m shoreline. However, contraction of the lake was accompanied by the development of the modern Okavango Delta. Timing of the contraction of the lake and initiation of the Okavango Delta is not tightly constrained, but by the time we propose that modern humans emerge within the region, some 200 ka, we speculate that the former extensive Makgadikgadi palaeo-lake had contracted to a much less extensive deltaic-lacustrine system, though nevertheless a favourable habitat for hominid occupation. It is this palaeo-wetland region that we propose as the homeland for AMH's founder population.

Climate Model Simulations and palaeo-climate data. To place the coalescence time estimates of the L0 branch into a climatic context and to test the robustness of simulated hydroclimate responses in South Africa to orbital-scale conditions, we use the LOVECLIM

earth system model of intermediate complexity²⁸. It is based on a 3-layer atmosphere, a 20-level ocean general circulation model, a dynamic-thermodynamic sea-ice model and a terrestrial vegetation model. A transient simulation covering several glacial/interglacial cycles was conducted by using time-dependent boundary conditions. The experiment²⁸ (covering past 784 ka) uses time-varying boundary conditions for orbital parameters, CO₂ and other greenhouse gas concentrations obtained from Antarctic ice cores, and an estimate of northern hemispheric ice-sheet orography and albedo changes (data are used in **Fig. 3** and Extended Data Fig. 6). The forcings are applied with an acceleration factor of 5: one coupled model year corresponds to five orbital calendar years. Our analysis focuses on the past 250 ka in both simulations. The climate sensitivity of this model to CO₂ variations was modified to capture the range of reconstructed global mean surface temperature changes in response to radiative forcing⁶². The transient LOVECLIM model simulations has previously been validated against other palaeo-climate records from around the world^{28,62,63}. Our analysis here focuses on the simulated precipitation as well as tree and grass fraction changes in central eastern Africa and western southern Africa (data used in **Fig. 3 d-f** and Extended Data Fig. 6 b,d).

As a result of its coarse horizontal atmospheric resolution (5.6°) and the use of only parameterized ageostrophic wind components, LOVECLIM has several deficiencies. Among the most noteworthy are the lack of realistic El Nino-Southern Oscillation variability and the fact that annual mean freshwater flux corrections have been applied to mimic the atmospheric moisture transport from the Atlantic to the Pacific and to stabilize the Atlantic Meridional Overturning Circulation.

There exist only a few long-term hydroclimate datasets from Southern Africa that cover the last >120-ka. Here we compare the simulated LOVECLIM precipitation (normalized) in central southern Africa with a southern central African hydroclimate composite, obtained by averaging the normalized orbitally-tuned rainfall reconstruction from the Pretoria salt pan²⁹ and

the normalized Fe/K river runoff proxy obtained from marine sediment core CD154-1006P³¹ (**Fig. 4b**). The composite index emphasizes the joint variability in both records. We find that some of the overall features in the observations, in particular the fact that rainfall is modulated by the precessional cycle of austral summer insolation⁶⁴ (**Fig. 3a**), are well captured by the LOVECLIM model experiment. However, we also find some discrepancies in the central part of southern Africa, such as in the phase of the precessional signal and the difference in overall wet and dry conditions during the Homeland period from 200-120 ka. The overall glacial drying in the central part of southern Africa from 100 ka to 20 ka is, however, captured in both, model simulation and palaeo-proxy reconstructions (**Fig. 3b,e**). Orbital-scale hydroclimate variations in southern Africa are clearly not spatially homogenous (**Fig. 3b-f**). To get a better understanding of the spatial patterns of hydroclimate variability we compared the model simulation with a composite index from southwestern Africa, obtained by averaging a normalized aridity index reconstructed from sediment core MD96-2094³³ and the normalized $\delta^{13}\text{C}$ isotope ratio data of leaf waxes extracted from the South Atlantic sediment core MD08-3167³² (**Fig. 3c** and Extended Data Fig. 6 c,d). The results show a good correspondence between model and reconstructions on the western side of southern Africa, and in particular reproduce a major drought period peaking ~120 ka and a subsequent increase of rainfall towards the last glacial period. This gradual increase in humidity/rainfall corresponds to an overall increase in lineage-splitting of the L0d1'2 haplogroup (**Fig. 3f**) and growth of its population (**Fig. 3c**). This result further highlights the possibility that climate shifts may have played an important role in the southwestward migration of L0d1'2 descendants (**Fig. 2**).

To further test the fidelity of LOVECLIM in reproducing interhemispheric orbital rainfall shifts across Africa, we also compared the simulated vegetation changes with a leaf-wax index from stable hydrogen isotope data extracted from a sediment core in the Gulf of Aden³⁰, which is indicative of hydroclimate and vegetation changes in the north-eastern Horn of Africa

(Extended Data Fig. 6b). The comparison shows a good qualitative correspondence for the precessional-scale timing of rainfall/vegetation maxima and minima as well as of the eccentricity modulated amplitude of these changes, lending further support to the credibility of the simulated rainfall patterns across Africa. It should be noted that regional patterns of paleo rainfall changes are in general difficult to simulate. In response to Last Glacial Maximum boundary conditions different Coupled General Circulation models simulate widely varying responses in rainfall over Africa²⁸.

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Extended Data

Extended Data Table 1. L0-mitogenomes included in this study. Footnotes: * Our previously published data; † Mitochondrial sequences of the coding-region only; ‡ Sequence has non-canonical start position corresponding to position 577 of rCRS; § Coriell cell lines.

Extended Data Table 2. KhoeSan population identifiers used in this study.

Extended Data Fig. 1. Phylogenetic tree of all 1,217 L0-mitogenomes. Phylogeny was inferred using FastTree v2.1.746, displayed using FigTree. Tips belonging to the same haplogroup are collapsed and coloured as in Fig. 2a. Local support values for each node are indicated and branch lengths are proportional to the number of substitutions per site. The tree is rooted to the seven Neanderthal mitogenomes indicated.

Extended Data Fig. 2. Detailed phylogenetic branching of L0k, L0d3, L0f, and L0g. Shown are the expanded sections of the phylogenetic tree depicted in **Fig. 2a**, including: **a**, 34 (of 113 total) L0k, **b**, all 40 L0d3, **c**, all 27 L0f, and **d**, all nine L0g mitogenomes. Each mitogenome is represented as a tip and coloured based on their broad ethno-linguistic classification, if known: KhoeSan in orange, non-KhoeSan in grey, and Cape multi-ethnic (KhoeSan-ancestral) in green. Public mitogenomes for which we cannot be certain of their broad population identifier are labelled in black font. Proposed new sub-lineages for L0d3, L0f, and L0g1 are indicated by red coloured node labels and further detailed in Supplementary Tables 7, 8, and 9.

773

774 **Extended Data Fig. 3. Detailed phylogenetic branching of L0d2.** Shown are the expanded
775 branches of the phylogenetic tree depicted in Fig. 2a, including: **a**, 51 (of 118 total) L0d2a, **c**,
776 25 (of 53) L0d2c, and **d**, all 11 L0d2d mitogenomes. For L0d2b, an additional BEAST analysis
777 was performed using an alternate subset of 441 mitogenomes that included all 43 L0d2b
778 samples, as opposed to the n=461 subset (Fig. 2a) that included only 13 L0d2b. The same
779 model parameters were used for both data subsets. In all panels, each mitogenome is
780 represented as a tip and coloured based on their broad ethno-linguistic classification, as in
781 Extended Data Fig. 2. The previously defined L0d2c1c haplogroup, containing the Coastal
782 KhoeSan StHe skeleton⁶ and other newly proposed sub-lineages are indicated by red node
783 labels (Supplementary Tables 4, 5, and 6).

784

785 **Extended Data Fig. 4. Detailed phylogenetic branching of L0d1.** Shown are the expanded
786 branches of the phylogenetic tree depicted in Fig. 2a, including: **a**, 54 (of 91 total) L0d1a, **b**, 45
787 (of 174) L0d1b, and **c**, 33 (of 184) L0d1c mitogenomes. Each mitogenome is represented as
788 tips and coloured here based on their broad ethno-linguistic classification as in Extended Data
789 Fig. 2.

790

791 **Extended Data Fig. 5. Detailed phylogenetic branching of L0a.** Shown is the L0a branch of
792 the phylogenetic tree displayed in Fig. 2a, which includes a subset of 114 (of 294 total) L0a
793 mitogenomes. Each mitogenome is represented as tips and coloured here based on their broad
794 ethno-linguistic classification as in Extended Data Fig. 2.

795

796 **Extended Data Fig. 6. Paleo-data/model comparison.** **a**, Relative rainfall changes (%)
797 between Eemian (125 ka) and pre-industrial conditions simulated by the Community Earth
798 System model using 20 km atmosphere and 10 km ocean resolution (see Methods). The
799 simulated rainfall differences are mostly due to enhanced (reduced) northern (southern)
800 hemisphere summer insolation. Locations of key sites that are used for paleo model/data
801 comparison in this study are highlighted in red. **b**, Simulated tree fraction (%) at Horn of
802 Africa (land grid points nearest to RC09-166) (grey, dark-blue bars) and stable hydrogen
803 isotopic composition of leaf waxes, corrected for ice volume contributions from the Gulf of
804 Aden marine sediment core RC09-166³¹, indicating changes in hydroclimate. **c**, Relative
805 precipitation changes (%) simulated by LOVECLIM transient model (all forcings) for 11°E,
806 19°S and grain-size aridity index reconstructed from sediment core MD96-2094³⁴. **d**, Grass
807 fraction changes simulated by LOVECLIM transient model (all forcings) at 11°E, 14°S-17°S
808 (dark blue, grey shading) and reconstructed $\delta^{13}\text{C}$ changes of n-alkanes (orange) (South Atlantic
809 sediment core MD08-3167) indicative of abundance of C_3 and C_4 plants in Namibia desert and
810 further inland³³.
811